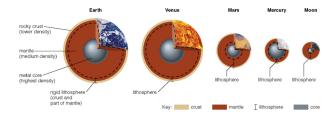


Interiors of the terrestrial planets

- Differentiated: dense, metallic core; inner part solid, outer fluid. lighter, mostly rocky mantle (*not* liquid, but can flow) even lighter rocky crust.

- Heat sources: accretion

- differentiation/contraction radioactive decay (other than residual heat) tidal
- Heat flow: conduction in core convection in (outer) mantle ("rock creep"; ~100 Myr timescale) conduction through rigid lithosphere



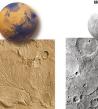
© 2005 Pearson Education, Inc., publishing as Addison Wesley

Exteriors of the terrestial planets

Shaped by four main processes:

- Impacts (many early on, now rare; can be used to "date" surfaces) Volcanism (Earth, Venus active; Mars maybe, maybe not; Mercury inactive)
- Plate tectonics (still active on Earth, not on Venus(?), Mars, Mercury)
- Erosion (wind, water, ...) (active on Earth, Mars; not on Venus, Mercury)

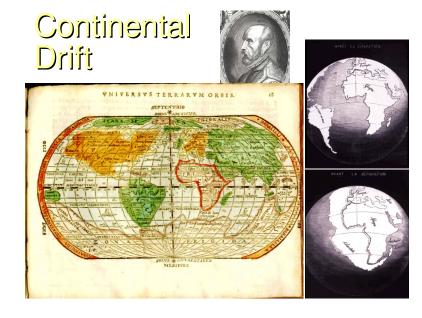


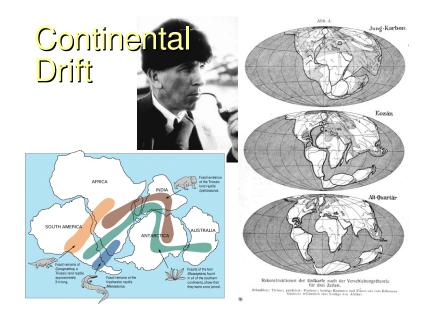


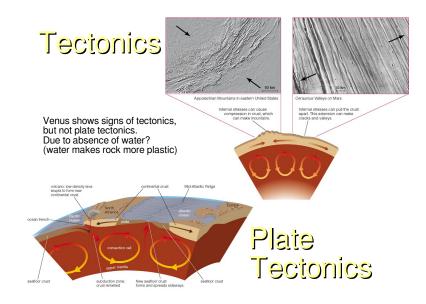
Earth has a variety of geological features visible in this photo from orbit.

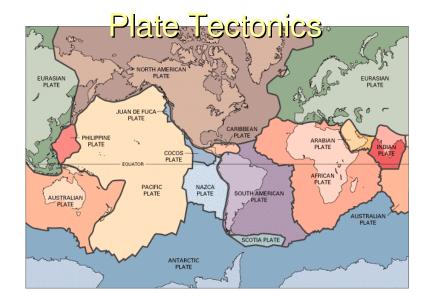
Mars has impact craters like the one The central structure is a tall, twin-peaked volcano on Venus near the upper right, but it also has features that look much like dried up

Mercury is beavily cratered, but also has long, steep clifs—one is visib here as the long curve that passes through the center of the image.

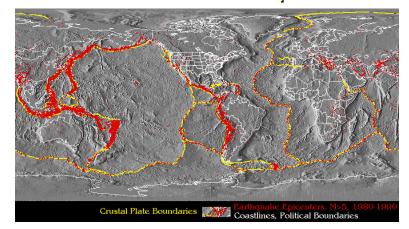




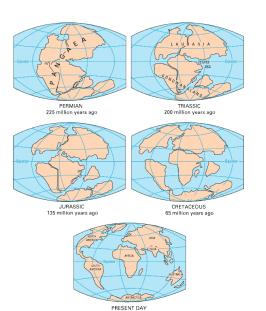




Plates and Earthquakes







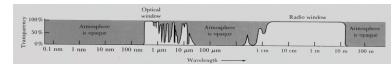
Passively Heated by the Sun --- the further the cooler

Typically we observe objects in reflected light, however, all objects emit re-processed thermal radiation which is observable at longer wavelengths.

Blackbody temperature for a non-self-luminous spherical body at distance ${f a}$ away from the Sun (with albedo A -- reflectivity)

$L_{abs} = (1-A) \frac{\pi R_{p}^{2}}{4\pi a^{2}} 4\pi R_{s}^{2} \sigma T_{s}^{4}; L_{em} = 4\pi R_{p}^{2} \sigma T_{p}^{4}$						
	$L_{abs} = L_{em}$, then		$T_{p} = \left(\frac{R_{o}}{2a}\right)^{1/2} T_{s} (1-A)^{1/4}$.) ^{1/4}	
	a (AU)	Α		T _{pred} (K)	T _{act} (K)	
Mercury Venus Earth Mars	0.4 0.7 1 1.5	0. 0.	06 77 30 25	422 K 230K 255K 218K	100-725 733 288 223	(?) (?) (?) good

Greenhouse effect: optical radiation from the Sun reaches the ground, but the infrared radiation from the ground cannot easily escape.



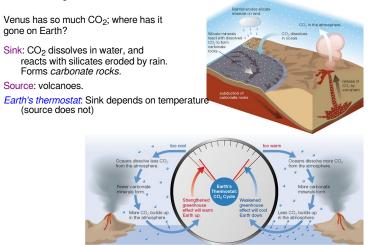
Atmospheres: Terrestrial Planets

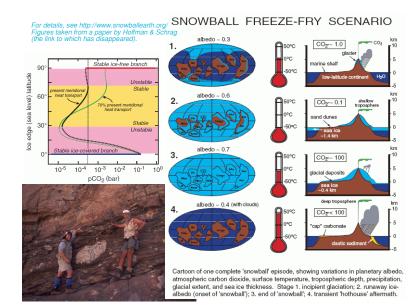
	Composition	surface pre	essure/T	
Mercury Venus Earth Mars	97% CO ₂ , 3% N ₂ 78% N ₂ , 21% O ₂ ,1% Ar 95% CO ₂ ,3% N ₂ ,1.6% Ar	< 10 ⁻¹² ba 92 bar 1 bar 0.006 b		(460°C) (15°C)
Titan (@S)95% N ₂ , few% CH ₄ ,Ar	1.5 bar	93 K (-	180ºC)
	spheres are reasonably we Ilar weight separation)	ell-mixed	Ar Carbon dioxide	ngon All others Oxygen
			Nitrogen	

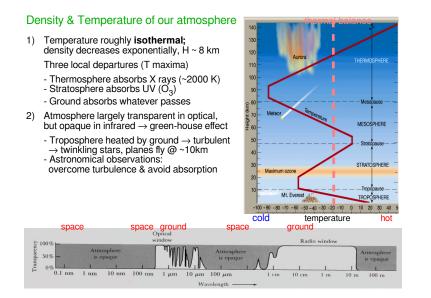
Earth's atmospheric composition

From http://www.ux1.eiu.edu/~cfjps/1400/atmos_origin.html

CO₂ cycle and Earth's Thermostat







Atmospheric optics: I) Why is the sky blue on Earth? Rayleigh scattering

air molecules & other constituents (N₂, O₂, H₂O droplets, dust...) all have sizes smaller than optical λ , and they preferentially scatter short- λ photons: $\sigma \sim 1/\lambda^4$

Earth: sky is blue (--> ocean blue) sunset is red (reddened) horizon whiter than zenith Fall/Winter sky dark blue UV is diffuse

Moon: sky is black

Mars: sky is reddish yellow

fine-dust (1-10µm) Mie scattering --> white iron oxide mineral absorption in the blue --> reddish

Mars Pathfinder true-color picture of Martian noon





Example of Rayleigh Scattering: Interstellar Reddening

- 1. interstellar space not empty
- 2. interstellar molecules & dust grains r < λ
- 3. scattered away blue; transmitted red

blue reflection nebula of Pleiades



Barnard 68

why is the Moon red during a lunar eclipse?



Atmospheric optics: II) Clouds

What are clouds?

How do they form?

Aggregates of water or ice droplets suspended in air In troposphere: low clouds-- water; high clouds-- ice 100% hum. + condensation nuclei (dust, cosmic-rays)

e.g., rising air that cools (--> humidity increases)

Why are clouds white? Water droplet colorless, solar light white

Mie scattering (droplets size $r \sim 10 \mu m > \lambda$), nearly geometric optics, no λ dependence (at sunset, cloud is red) soap foam: geometric scattering, also no λ dep.



Why don't clouds fall from the sky? Tiny droplets, fall slowly; updraft mixing? Fall and evaporate and form new ones? Electrically charged clouds?



Origin of Earth's atmosphere

Our (& Venusian) atmosphere cannot be primordial

- 1) N₂, CO₂, H₂O are not condensed at 1AU from Sun, O₂ does not naturally occur
- 2) Earth too low in mass to accrete gas directly
- Gas is unlikely to have been trapped in solids and dragged to Earth, since noble gases (Ne, Kr, Xe) are heavily depleted relative to solar abundance.
- New-born Earth molten and hot (10³K) --> most gases can escape thermally.

Some relief only in that in the early bombardment period (~ 700 Myr) water can be brought in by comets & asteroids. (*Note: D/H ratio in comets ~2 higher than ocean, so these cannot do it alone*)

Origin of Earth's atmosphere (cont'd)

Our atmosphere is obtained gradually: volcanic outgassing & invaders

1 st atmosphere thermal escape H & He(?)	2 nd atmosphere outgassing/accretion CO ₂ /NH ₃ outgassed H ₂ O accreted/outgassed	3^{rd} atmosphere absorbing CO ₂ most H ₂ O liquid CO ₂ got locked in	
P: ? T: ~10³K	(solid crust/ocean, 3.5Gyrs ago) ~100 bar (like Venus!) 0°C< T < 100°C	O ₂ produced ~1 bar ~15°C	

 sinks of CO2:
 sedimentary rock via H2O, life (carbon) via photon-synthesis

 sources of CO2:
 volcanic outgassing (+human activities)

 sinks of H2O:
 subducting plates

 sources of H2O:
 outgassing, comets/asteroids?

Currently sensitive balance reached, mild green-house

run-away green-house: too much CO₂, H₂O can all disappear

 \rightarrow sink disappears as well while outgassing produces yet more ${\rm CO_2}$

Venus: divergent evolution from Earth

	a(AU)	mass(M _F)	spin atm.	Pressure	Т	plate tect.	ocean
Earth:	1	1 -	1 day	1bar	288 K	Yes	Yes
Venus:	0.7	0.8	243 day	92bar	770 K	No	No

1) 97% CO₂ in the atmosphere, ~ 700K, *no CO₂ sink due to dryness*

2) Why so dry? high D/H ratio indicates past large H₂O reserve Green-house runaway and H₂O photo-evaporated

3) Cratering no older than ~0.8 Gyr \rightarrow tectonics stopped recently

A planet is a nonlinear system. Strongly divergent evolution can occur.

Cause & Effect?

1) Slightly closer to the Sun and got torched? Or formation site had naturally less H₂O?

2) Too much CO₂ to start with and H₂O never condensed

(But: Initial Earth atm. ~100 bar, mostly CO2 --> would require fine tuning?)

The Story for Mars: 2nd atmosphere gradually lost, no outgassing (tectonics)



Origin of O₂ on Earth: photosynthesis; $CO_2 + H_2O + hv \rightarrow O_2 + carbo-hydrate$

